

59°45'N

Mike Gracz 907-299-6122

151°37'30"W

59°45'N

Temperatures in exceedance for returning adult salmon



Figure 20. Temperature profiles of sites in the northern portion of the Anchor River watershed.

More Flow = Temperature Buffer

More Flow = Temperature Buffer

Peatlands important for dry-season flow?

"The idea that intact mires act as useful water storage reservoirs...is clearly wrong: drained mires are better reservoirs and mineral soil ecosystems are better still." -H.A.P. Ingram 1983.

"...we now know that many bog peats *do not* typically act like 'sponges'...Rather, baseflows are poorly maintained..." -J Holden et al 2004 "The idea that intact mires act as useful water storage reservoirs...is clearly wrong: drained mires are better reservoirs and mineral soil ecosystems are better still." -H.A.P. Ingram 1983.

"...we now know that many bog peats do not typically act like 'sponges'...Rather, baseflows are poorly maintained..." -J Holden et al 2004

"During the drought, streams draining blanket peat on Plynlimon and elsewhere In mid-Wales sustained higher minimum flows than those draining podzol or brown-earth soils."

-M.D. Newson 1980

"Thus, more than half of the...runoff... during years of near and below normal precipitation... is water slowly released from storage within lakes and wetlands." -E.A. Ackroyd et al. 1967

Hydrogeologic setting is important

0.2 Kilometers



and a second second second





Water produced during drawdown?



Water Produced during drawdown?



= Specific Yield (S_v)



Potential Evapotranspiration (ET): Thornthwaite method Temperature Latitude

Diurnal Method

Night-time rate of decline



Diurnal Method

Day-time rate of decline



Diurnal Method

Difference = ET



Surplus Remaining For Stream Flow Limpopo Creek

Actual water level decline (228mm/8d) Area of similar peatlands in watershed

Actual dry-period flow (0.06 m³s⁻¹)

Amount Remaining For Stream Flow (Actual dry-period flow = $0.06 \text{ m}^3\text{s}^{-1}$)

	Specific Yield			
ф	0.45	.14	.10	0.05
0.8	0.259	0.080	0.060	0.029
0.9	0.297	0.092	0.066	0.033
0.131	0	0	0	0
Diurnal ET	0.108	0.033	0.024	0.012

Possible that peatlands support flow during dry periods

% contribution of end-members to stream flow











Mixing Analysis



Fig. 3. Artificial data with pure end-members in U space defined by the correlation matrix.

FROM:

Christophersen N, Hooper RP 1992. Multivariate analysis of stream water chemical data: The use of principal components analysis for the end-member mixing problem. *Water Resources Research* 28, 99-107. DOI: 10.1029/91wr02518.

We have a problem.....

A funny thing happened on the way to the stream....



Flow & Rainfall average over time Chloride in rain averages over time



Amount of rainfall or streamflow Chloride in rainfall or stream water

Flow averages over time Chloride in rain averages over time Chloride in stream water does not average



Amount of rainfall or streamflow Chloride in rainfall or stream water



Fig. 4. Non-self-averaging behavior in water quality time series, illustrated by rms differences between successive mean concentrations of selected solutes in 7-h and weekly samples of Upper Hafren streamwater (solid and open symbols, respectively) averaged over intervals ranging from 7 h to 5–10 y. Error bars show SEs. Thin gray reference lines show trends for non-self-averaging behavior, in which averages over longer and longer time scales do not converge. Heavy gray lines show the slope of -0.5 predicted by the central limit theorem for self-averaging time series. The solutes generally plot as horizontal lines, indicating non-self-averaging behavior. In contrast, stream discharge and its logarithmic transform both follow the self-averaging behavior indicated by the heavy gray lines, for time scales longer than ~ 0.1 y. Individual solutes are shifted by arbitrary factors so they can be plotted together. Plots for all 45 solutes and both sampling sites are shown in *Sl Appendix*, Fig. S10.

NOT Self-averaging over time: Lead (or almost everything but streamflow)

Self-averaging over time: streamflow

From: Kirchner, J.W., and Neal, C., 2013. Universal fractal scaling in stream chemistry and its implications for solute transport and water quality trend detection. Proceedings of the National Academy of Sciences 110, 122213-122218

Watershed acts as a "fractal filter"



Mixing Analysis





Water samples analyzed: Cations on ICP-MS at UAA-ASET lab (B. Hagedorn) Isotopes at ENRI Stable Isotope Lab (J. Welker) Anions at EPA MED lab, Duluth, MN (M. Moffett, L. Anderson)





Water samples analyzed: Cations on ICP-MS at UAA-ASET lab (B. Hagedorn) oppes at ENRI Stable Isotope Lak (I. Welker) Anions at EPA MED lab, Duluth, MN (M. Moffett, L. Anderson)





-Hooper 2003

DL=0.01

\$04

0,8194

0.3452

-0.30883 0.496408 -0.28425 -0.47652 -0.51367 0.131514 0.154959 0.160938 -0.1192 -0.20744 0.199012 -0.75956 0.135297 0.693666 0.84249 -0.4021 0.995627 0.476578 0.27173 138849 0 0.91819 1.526556 0.999845 0.76136 0.8561 -0.07828 -0.94632 -0.04216 0.26794 1.83777 2.511386 -1.73969 -2.44647 -3.5418 1 278723 -0.54661 1.341233 0.646405 0.371684 -0.23745 0.266002 -0.23114 -0.26467 -0.26035 -0.60898 -0.40909 -4.3168 0.653961 -1.32022 0.572613 1.245622 1.400724 0.216455 -0.92188 -0.30401 -0.0757 dex shat* (index allows manipulation of matrix eleme

X*VE(VVE)

-0.30883 0.496408 -0.28425 -0.47652 -0.51367 0.131514 0.154959 0.160938 -0.1192 -0.20744 1 949961 -0.4021 0.995627 0.476578 0.2717 1.494836 0.91819 1.526556 0.999845 0.76136 -0.8561 -0.07828 -0.94682 -0.04216 0.26794 -1.83777 2.511386 -1.73969 -2.44647 -2.54185 1.278723 -0.54661 1.341233 0.646405 0.371684 0.23745 0.266002 -0.23114 -0.26467 -0.26035 .377749 -0.60898 -0.40909 -0.31683 n 653961 -1.32022 0.572613 1.245622 1.400724 0 86329 0.218455 -0.92188 -0.30401 -0.0757 0.01398 15.24685 0.043498 0.19385 -0.0141 19.60222 0.047168 0.21208 -0.01441 19.35136 0.055517 0.274445 22 20000 -0.01429 27.76817 0.053287 0.240543

0.01446 32.96237 0.058661 0.269625

-0.01329 1.007924 0.023266 0.078423

13.8215 -0.01418 8.769612 0.047959 0.240318



Peat sample collection

5 tracers: SO₄, K, δ¹⁸O, Ni, Ba

3 End-member model

55% Matches peatland flow in water budget

	Specific Yield				
ф	0.45	.14	0.10	0.05	
vaite 8.0	0.259	0.080	0.060	0.029	
0.9	0.297	0.092	0.066	0.033	
Ĕ Ĕ 0.131	0	0	0	0	
Diurnal ET	0.108	0.033	0.024	0.012	

(Actual streamflow = $0.06 \text{ m}^3\text{s}^{-1}$)

3 End-member model

Low conductivity of Glacial Till

(Stream length x width)(k_{till}) = Flow_{till}

 $14,438 \text{ m x } 2 \text{ m x } 10^{-8} \text{ m} \cdot \text{s}^{-1} = 0.0029 \text{ m}^3 \text{s}^{-1}$

= 4.8% of low stream flow

(4.5% from EMMA)

1. ET estimated using a diurnal method without recharge

2. Space-for-Time sampling in EMMA can avoid fractal filtering problems

3. Peat contributed ~55% of dry-season stream flow

Special Thanks:

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